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GEOLOGICAL INVESTIGATION OF THE MISSISSIPPI RIVER AREA ARTONISH TO DONALDSONVILLE, LA.

by

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GEOGRAPHIC SETTING

- 1. The area under consideration in this report encompasses some 2500 square miles that flank the Mississippi River between a point about 4 miles north of Artonish, Louisiana, and a point about 15 miles east of Donaldsonville, Louisiana (fig. 1). It lies almost entirely within the New Orleans District of the Corps of Engineers, extending from Mississippi River miles 152 to 327 above Head of Passes. Baton Rouge, Louisiana, is by far the largest urban complex within the area.
- 2. Physiographically, the study area includes the extreme southeastern part of the Lower Mississippi Valley or the Mississippi River floodplain plus a smaller area of terracelands or bluff-lands of Pleistocene age lying east of the floodplain. Throughout the area, the Mississippi River flows close to the base of the bluffs, which decrease in height downstream until, at Donaldsonville, they rise only as a low terrace 10 ft or less above floodplain level. The floodplain area west of the river is part of the large Atchafalaya Basin of Louisiana. Toward the south and southeast, the study area merges with the Mississippi River Deltaic Plain.

PURPOSES AND SCOPE

3. The purposes of this investigation have been to (a) determine the areal distribution and physical characteristics of the various alluvial deposits, (b) analyze subsurface soils conditions of various environments of deposition to aid in determining foundation and underseepage conditions, (c) determine the nature of and the depth to the Tertiary or Pleistocene deposits lying beneath the Recent alluvium, and (d) delineate the major geologic formations or basic lithologic types in the uplands or blufflands. During the past decade, the U. S. Army Engineer Waterways Experiment Station (WES) has been conducting continuing investigations of the Yazoo, St. Francis, and Boeuf-Tensas Basins of the Lower Mississippi Valley with similar goals. In all cases, including the present study, the basic data presentation form is a standard 1:62,500-scale topographic quadrangle supplemented by one or more geologic cross sections. The 10 quadrangles included herein are thus part of what will ultimately be almost complete engineering-geologic quadrangle map coverage of the Lower Mississippi Valley area. A loose-leaf folder format has been adopted for the report series to facilitate addition of supplemental data and/or revisions.

MAPPING PROCEDURE

- 4. The areal distribution of the Recent environments of deposition was determined largely from aerial photos and photo mosaics ranging in scale from 1:10,000 to 1:63,360 and in date from the 1930's to the 1960's. Pertinent geological publications such as groundwater investigations and state geological survey bulletins, and unpublished data from manuscript reports and graduate theses (see plates for references) provided much of the data on the Pleistocene and Tertiary formations exposed in the uplands and beneath the Recent alluvium.
- 5. A large volume of subsurface data was obtained from the files of agencies such as the U.S. Army Engineer District, New Orleans, the U.S. Geological Survey, the Louisiana Department of Highways, several private foundation engineering firms, and numerous private industries with plants in the area. The more detailed subsurface information, usually the logs of holes drilled by the Corps of Engineers, was used to construct cross sections through the various quadrangle areas and to contour the surface of the entrenched valley.

GEOLOGIC SETTING

- 6. Deposits of Tertiary age are exposed in the uplands only at a few small and scattered localities north of St. Francisville, Louisiana (fig. 1). These materials everywhere are fine-grained and are identified as undifferentiated Pliocene-Miocene deposits, probably being either the Pascagoula or the Hattiesburg formation equivalents (see reference 1 for further information on Tertiary stratigraphy). Fluvial deposits of Pleistocene age, characteristically graveliferous, form a nearly continuous thin blanket over the Tertiary formations north of Zachary, Louisiana (reference 2 contains detailed descriptions of the deposits). South of this point (fig. 1), the fluvial deposits are in the form of well-defined terraces that lack near-surface graveliferous units and that thicken rapidly southward to thicknesses of at least several hundred feet (reference 3). The youngest terrace, designated as the Prairie formation (Qtp) according to Fisk (reference 4), is widespread and easily identifiable. Although all of the older Pleistocene deposits north of Zachary have also been differentiated by Fisk into terrace formations, recent investigations by Parsons (reference 2) and others have indicated that alternate interpretations are tenable but not definitive. Such deposits are identified only as undifferentiated Pleistocene deposits (Qtu) in this report.
- 7. The base of the Recent alluvium is an entrenched surface displaying a dendritic drainage pattern. The average elevation of the surface is well over 100 ft below floodplain level, and maximum depths of as much as 350 ft below floodplain level occur in the deepest trench. The entrenched surface was created during late Pleistocene times as a result of the Mississippi River and various tributaries adjusting their gradients to lower-than-present sea levels during periods of continental glaciation. North of the latitude of Zachary, Louisiana, the entrenchment affected Tertiary formations, while south of this latitude, Pleistocene formations are present and form the entrenched surface. The contact between the two is interpreted to be along a major fault zone known as the Bancroft Fault Zone (reference 5). This fault zone (shown in plates Fordoche (a) and New Roads (a)) is one of several in the study area that trend roughly east-west and that are composed of parallel normal faults with near-surface displacements of several tens of feet.
- 8. The oldest Recent deposits filling the entrenched valley are included in the thick wedge of fluvial substratum sands and gravels. This unit is by far the thickest and most continuous body of sediments of essentially one type that occurs in the study area or in the entire Mississippi Alluvial Valley. At occasional points above the deeper entrenchments, substratum deposits attain a thickness of 300 ft. The typical substratum sequence is fine sands grading downward into progressively coarser sands. The first gravels appear at depths of 75 to 150 ft and become more abundant and larger in size with increasing depth.
- 9. Substratum deposits below a depth of 80 to 100 ft were deposited by shallow, swiftly flowing, braided courses of the Mississippi River when it was carrying large volumes of coarse glacial debris. Those substratum deposits above this depth, largely limited to the area of the present Mississippi River meander belt, were deposited by the river since it changed from a braided to a meandering regimen, possibly about 12,000 years ago.
- 10. Deposits laid down since that time are all relatively fine-grained and are subdivided according to environment of deposition. Figure 2 summarizes the characteristics of the six major environments recognized in the study area that collectively are referred to as topstratum deposits.
- 11. Sediments deposited in the backswamp environment are by far the thickest and areally most extensive in the study area. Initial deposition began about 12,000 years ago when the Mississippi River followed a course along the western side of its alluvial valley and continued uniformly across the study area until about 4,000 years ago, at which time the river shifted to a course along the eastern side of the valley. Backswamp deposits continued to accumulate along the flank of the new meander belt; however, meandering of the river within the new meander belt resulted in the erosion and re-

moval of large areas of backswamp deposits. These deposits were replaced largely with point bar deposits and deposits that accumulated in abandoned channel environments.

12. Several small distributaries such as Bayou Latenache, Bayou Fordoche-Bayou Grosse Tete, and Bayou Plaquemine (see plates for locations) are apparent in the area because of their well-developed natural levee ridges. These distributaries were apparently short-lived, carried only a small volume of flow only at flood stages on the river, and originated as crevasses. Bayou Lafourche is the only distributary in the study area that carried all or a significant portion of the Mississippi River discharge (reference 6).

DATA PRESENTATION

- 13. The plates in this folio show the distribution of alluvial deposits in plan and in profile. On each of the base maps (plates designated "a"), which are full-scale reproductions of the latest standard 1:62,500-scale topographic quadrangles, four of the six environments of deposition of the topstratum are shown in color. The other two environments, the alluvial apron and the natural levee deposits, are shown as a dashed and a dotted overprint, respectively, in order that the types of deposits lying beneath these two essentially surficial deposits will not be masked. Heavy black dashed lines are used to show the locations of selected major swales in point bar areas that illustrate the trends of the meanders.
- 14. A green line pattern is used to delineate that part of the backswamp environment where the deposits are unusually thin. This situation occurs along a narrow, discontinuous band that is interpreted as being the location of the meander belt of the local valley drainage system that was situated near the eastern valley wall while the Mississippi River was flowing near the western valley wall. The thin backswamp and natural levee deposits now overlying and obscuring all surface expression of the meander belt were deposited after the Mississippi River adopted its present course near the eastern valley wall.
- 15. A blue checkered pattern is used to portray large swalelike areas of various origins (e.g., crevasse scourings and areas of slack-water deposition behind bars) where thick deposits of soft, fine-grained materials (almost exclusively clays) occur. Certain areas of point bar environment are delineated by a fine dotted red overprint. In these areas, the point bar deposits are considerably finer grained than is typical for the environment. This situation develops where the river is forced to meander in an anomalous manner due to its impingement against the highly erosion-resistant Pleistocene or Tertiary uplands. Deposition of fine-grained materials (mostly silts with clay layers) actually results from an eddy and/or slack-water condition that develops in the bend.
- 16. The elevations of the surfaces of the entrenched Tertiary or Pleistocene formations are shown by red contours. The borings used to contour the surfaces are shown as small red dots.
- 17. Where boring information is sufficient, one or more cross sections have been prepared to accompany each map. Each plate containing cross sections bears the designation "b." Where information is sufficiently detailed, principally where closely spaced engineering borings have been made, the soil types are shown in color. Note that soil types are shown only to the depths of the detailed borings.
- 18. The classification of soil types used in the cross sections is based on the system used by the Lower Mississippi Valley Division, CE, prior to 1950. This was unfortunately necessary because of the large number of borings used in the study that predate 1950. For comparison with more recent borings classified by the Unified Soil Classification System (USCS), and in order that these borings could be used in the study, probable equivalents of the older system and the USCS were determined and are shown in the legend. It is emphasized, however, that the two systems cannot be equated precisely; for example, soils classified as lean clay (CL) according to the USCS may occasionally be included with the older system soil types shown in blue (clay sand, sandy clay, silt, sandy silt) as well as with those shown in green (clay, blue mud, silty clay, clay silt).

MAPPING LIMITATIONS

19. The maps and cross sections in this folio should be considered as being of a reconnaissance nature only. The aerial photo interpretation could be field checked to only a limited degree, and borings were not available as substantiating data in many areas. Furthermore, the mapping technique allows for little quality control, i.e. features of doubtful origin and/or areal extent necessarily are portrayed in the same manner as are well-defined ones of unequivocal origin. Also, it should be kept in mind that the accuracy of individual contacts, contours, and other designations is affected by the scale of the map, limitations in the source data, and progressive errors that may develop during the several stages of drafting and printing. In no case should an accuracy in plan of more than ±200 ft be expected for the position of a contact, contour, or other designation.

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TYPE OF **TOPSTRATUM** APPEARANCE ON AERIAL PHOTOGRAPHS METHOD OF DEPOSITION OCCURRENCE AND CHARACTERISTICS DIAGRAMMATIC ILLUSTRATION **DEPOSITS** Natural levees are low ridges which flank both The largest and most widespread natural levsides of streams that periodically overflow their ees in the area occur along the present course banks. Since the coarsest and greatest quantiand major abandoned distributaries of the Misties of sediment are deposited closest to the sissippi River. They attain crest heights of 15 to stream channels, the natural levees are highest 20 ft above the adjacent backswamp areas, and and thickest in these areas and gradually thin may be 3 miles or more in width. Natural levees away from the channels. In general, the greater NATURAL also occur along smaller streams in the area; the distance from the stream, the greater the however, they are appreciably narrower and LEVEE percentage of the finer grained sediments. Misteeper than those along the Mississippi River. nute drainage channels trending at right angles Typical natural levee deposits consist of stiff to the parent stream (down the backslope of the to very stiff, brown to grayish-brown silts, silty levees) are rather common; major crevasses clays, and clays that exhibit moderate to high are indicated when these channels are large and degrees of oxidation. Natural water contents of pronounced. Abandoned crevasse channels are the deposits are typically low, and organic matter often filled with sediments that are distinctly is seldom present except in the form of roots. coarser than the remainder of the natural levee. Alluvial apron deposits are not widespread in this area: they occur intermittently along Alluvial aprons are combinations of alluvial the eastern valley wall north of St. Francisville, and colluvial deposits which overlie the flood-La. They are best developed near the mouths of plain deposits along the valley walls and along the small streams that enter from the uplands the sides of upland remnants within the valley. and particularly where they overlie backswamp Typically, symmetrical alluvial fans are presdeposits and thus have not been affected by river ent at the mouths of streams that drain the uplands. When these streams are rather closely Reflecting the composition of the materials in the uplands (loess and terrace deposits primar-ALLUVIAL spaced, the fans coalesce to form the alluvial APRON aprons. When the streams are more widely ily), the alluvial apron deposits consist of clayey spaced, the fans are separated, and the intersilts, silts, and fine sands. Coarser sand and vening portions of the aprons are composed possibly even small quantities of gravel may be present near the mouths of the more active upland mainly of sediments that have washed down from the uplands or that have moved downslope by streams. Because they are well drained, they soil creep (colluvial deposits). are oxidized and generally similar to natural levee deposits. Point bar deposits consist of sediments laid down on the insides of river bends as a result Point bar deposits are widespread only along of meandering of the stream. Although the dethe present course of the Mississippi River, posits extend to a depth equal to the deepest porparticularly north of Baton Rouge, La. The tion or thalweg of the parent stream, only the topstratum deposits consist of tan to gray clays, uppermost, fine-grained portion is included as clayey silts, silts, and fine sands in the ridges, part of the topstratum. Within the point bar topand soft gray clays and silty clays in the swales. Typically, the point bar topstratum (excluding stratum, there are two types of deposits: silty and sandy, elongate bar deposits or "ridges" major swales) averages about 20 to 40 ft thick. POINT BAR which are laid down during high stages on the However, in certain areas where the river has stream, and silty and clayey deposits in arcuate meandered anomalously due to impingement depressions or "swales" which are laid down against the valley wall, the topstratum may be as during falling river stages. Characteristically, much as 80 to 100 ft thick. Both water and organic the ridges and swales form an alternating secontents are high in swales and in the anomalously ries, the configuration of which conforms to the thick topstratum areas, whereas they are both curvature of the migrating channel and indicates commonly low in the ridge deposits. the direction and extent of meandering. Backswamp deposits are by far the most characteristic materials in the area, being present and virtually uninterrupted by other deposits over Backswamp deposits consist of fine-grained thousands of square miles. Average deposit sediments laid down in broad, shallow basins durthicknesses vary from about 75 to 80 ft in the ing periods of stream flooding. The sedimentnorthern part of the area to about 110 ft in the carrying floodwater may be ponded between the southern part. Total thicknesses of fine-grained natural levee ridges on separate meander belts, materials (topstratum deposits) may reach 130 ft or between natural levee ridges and the uplands where natural levee deposits overlie backswamp **BACKSWAMP** or upland remnants within the alluvial valley. deposits. Backswamp areas typically have very low relief Soft to stiff, gray to dark gray-brown clays and and a distinctive, complicated drainage pattern silty clays are the typical backswamp deposits. in which the channels alternately serve as trib-Occasional thin layers of silt or sand may be utaries and distributaries at different times of found, and organic matter in the form of dissemthe annual flood cycle. inated particles, peat layers, and large wood fragments is numerous. Average water contents of the deposits are moderately high but less than those of channel and swale fillings. Abandoned channels, or "clay plugs" as they Abandoned channels of the Mississippi River are commonly called, are partially or wholly are not numerous in this area and occur only filled segments of stream channels formed when north of Baton Rouge, La. Individual abandoned the stream shortens its course. Soon after forchannels vary in length from a few miles to mation, they are usually characterized by open 20 miles or more, have widths of 3000 to 5000 ft water or oxbow lakes. Subsequently, they may or more, and usually exceed 100 ft in depth. become essentially filled and occasionally com-The upper portions of the arms of the loops of **ABANDONED** pletely obscured by various meander belt deneck cutoffs are normally filled with a short posits. The abandoned segment may represent CHANNEL wedge of fine sand and silty sand. The soft, gray an entire meander loop formed by the stream or blue-gray clays with high water contents that cutting directly across a narrow neck of two occur around the loop between the sand wedges converging arms of a loop (a neck cutoff), or it comprise the "clay plug" portion of the abanmay represent a portion of a loop formed when doned channel. Homogeneous, soft, fat clays 100 a stream occupies a large point bar swale durto 120 ft thick have been encountered in Missising flood stage and abandons the outer portion sippi River clay plugs. of the loop (a chute cutoff). Abandoned courses are lengthy segments of a river abandoned when the stream forms a new Bayou Lafourche marks the position of the only course across the floodplain. The abandoned major Mississippi Riverabandoned distributary course, varying from a few miles up to hundreds in the area. There are no major abandoned of miles in length, gradually fills with sediment courses in the area mapped and all other distriband is often occupied by a smaller or underfit utaries are small and were apparently quite short stream. In many cases, the smaller stream me-**ABANDONED** anders within the confines of the larger meander Only sparce data are available on the nature of **COURSE** belt and destroys segments of the abandoned the deposits filling the abandoned distributaries. course. In other cases, the smaller stream de-AND A large downstream-thinning sand wedge is lineates the extent of the abandoned course when known to be present in the Lafourche distributary; DISTRIBUTARY there are no other indications of its presence. however, the smaller distributaries appear to be Abandoned distributaries, characteristic feafilled almost entirely with fine-grained sedi-

tures of the deltaic plain, are analogous in both

morphology and lithology to abandoned courses

although any one distributary may never have

carriedall or even a substantial part of the total

stream discharge.

ments. The physical characteristics of the fine-

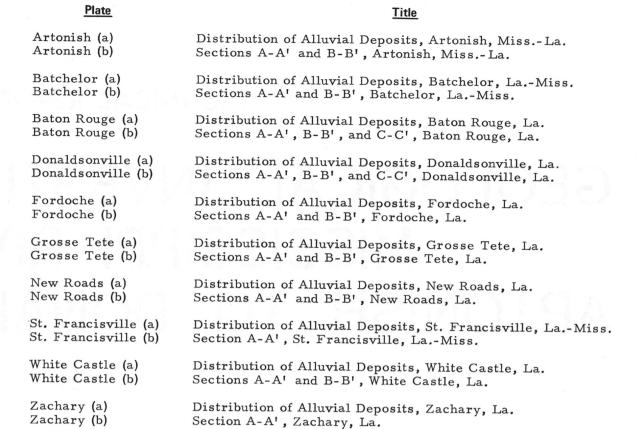
grained materials should be essentially the same

as those of abandoned channel deposits.

Authorization for this study is contained in a letter from the Division Engineer, U. S. Army Engineer Division, Lower Mississippi Valley, to the Director, U. S. Army Engineer Waterways Experiment Station (WES), dated 9 July 1965, subject, "Status of Soils Division Projects for MRC and LMVD for FY 1965 and Request for Funds for Projects for FY 1966."

The collection and interpretation of data for this study and the preparation of the text and plates for this report were accomplished by Dr. R. T. Saucier, Geology Branch, Soils Division, WES. All work was conducted under the direct supervision of Dr. C. R. Kolb, Chief of the Geology Branch, and Mr. W. B. Steinriede, Jr., Chief of the Civil Projects Section, Geology Branch, and under the general supervision of Messrs. W. J. Turnbull and A. A. Maxwell, Chief and Assistant Chief, respectively, of the Soils Division, WES.

Directors of the WES during the conduct of this study and the preparation of this report were COL John R. Oswalt, Jr., CE, and COL Levi A. Brown, CE. Technical Directors were Mr. J. B. Tiffany and Mr. F. R. Brown.



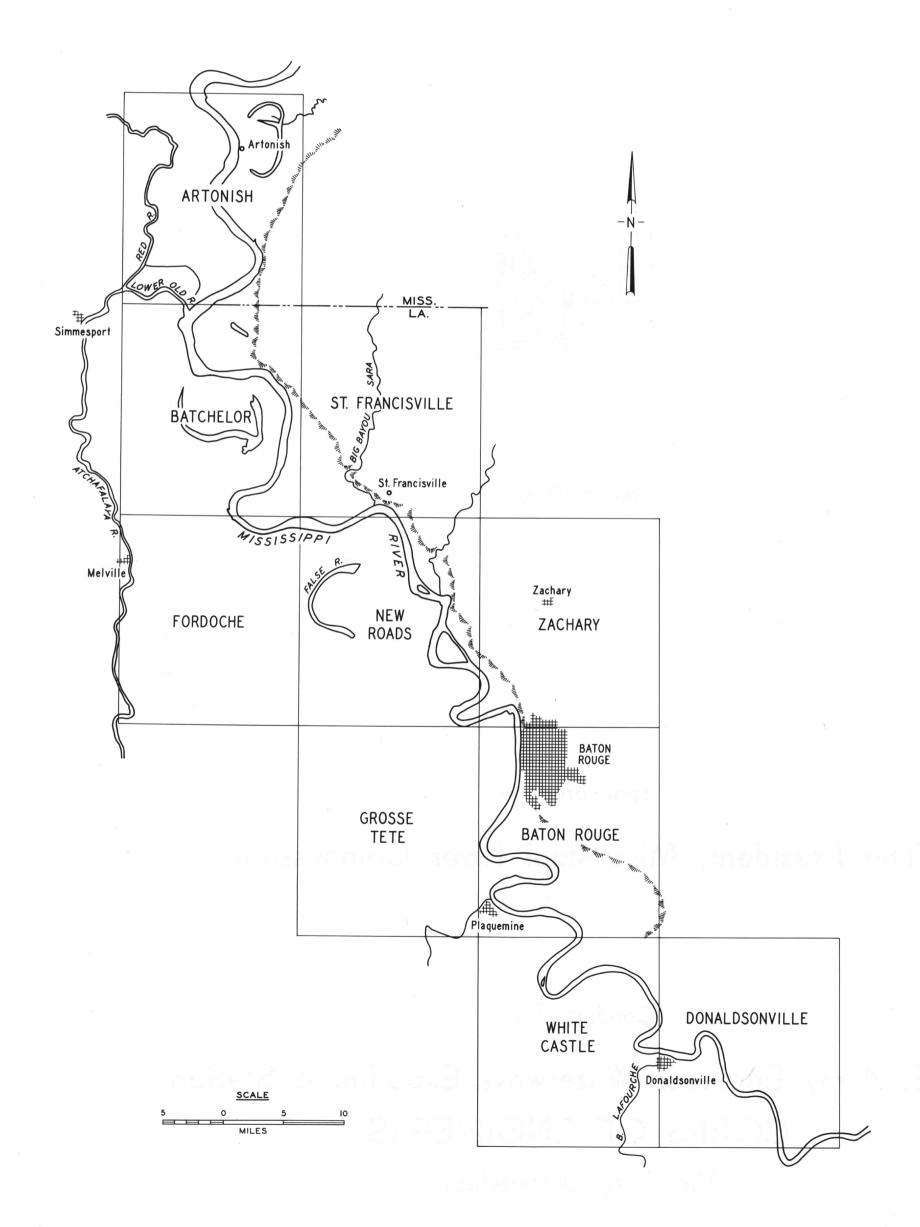


Fig. 1. Quadrangle coverage of the Mississippi River area, Artonish to Donaldsonville, La.